From Milankovitch to El Niño: Using proxies and models to understand the dynamics of the climate system

Steven J. Phipps ARC Centre of Excellence for Climate System Science Climate Change Research Centre University of New South Wales, Sydney, Australia

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Introduction

The "handshake" question



How do we integrate proxy data and climate models in a way that extracts the maximum possible information about the dynamics of the climate system?

< 17 ▶

The "handshake" question



- Data-model integration is a two-way process
- Proxy data can be used to constrain climate model simulations
- Climate models can provide dynamical interpretation of proxy data
- Everyone wins: we learn more about the dynamics of the climate system than when we employ the two approaches separately

ENSO over the past 8 ka

We know that ENSO has changed over the Holocene



Steven J. Phipps, ARC CoE for Climate System Science and Climate Change Research Centre, UNSW, Sydney, Australia

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Orbital changes have been the dominant external signal ...



... causing large seasonal changes in insolation



Climate model simulations

• The CSIRO Mk3L climate system model (Phipps et al., 2011, 2012)

- Atmospheric general circulation model $(5.6^{\circ} \times 3.2^{\circ}, 18 \text{ levels})$
- Ocean general circulation model $(2.8^{\circ} \times 1.6^{\circ}, 21 \text{ levels})$
- Dynamic-thermodynamic sea ice model
- Land surface scheme
- 10,000-year pre-industrial control simulation
- Three transient simulations of the past 8 ka:
 - Only the Earth's orbital geometry is varied (Berger et al., 1978)
 - Atmospheric CO₂ concentration = 280 ppm
 - Solar constant = $1365 \, \text{Wm}^{-2}$
 - Each ensemble member is initialised from different years of the control simulation (i.e. a perturbed initial conditions ensemble)

Simulated amplitude of ENSO variability (500-year mean)



Trend in August surface air temperature (K ka $^{-1}$)



Trend in August MSLP (hPa ka^{-1}) and surface wind stress



Trend in August precip (mm ka^{-1}) and surface wind stress



With weaker trades, El Niño events develop more easily



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ENSO over the past 1100 years

ENSO also changes on shorter timescales



Natural/anthropogenic forcings over the past 1100 years



Climate model simulations

• The CSIRO Mk3L climate system model (Phipps et al., 2011, 2012)

- Atmospheric general circulation model (5.6° × 3.2°, 18 levels)
- Ocean general circulation model $(2.8^{\circ} \times 1.6^{\circ}, 21 \text{ levels})$
- Dynamic-thermodynamic sea ice model
- Land surface scheme
- 10,000-year pre-industrial control simulation
- Three transient simulations of the past 1500 years:
 - Orbital changes (Berger, 1978)
 - Anthropogenic greenhouse gases (MacFarling Meure et al., 2006)
 - Solar irradiance (Steinhilber et al., 2009)
 - Explosive volcanism (Gao et al., 2008)
 - Each ensemble member is initialised from different years of the control simulation (i.e. a perturbed initial conditions ensemble)

Reconstructed/simulated ENSO amplitude (30-year mean)



ENSO amplitude versus individual forcings

Ensemble	Greenhouse	Solar	Volcanic
member	gases	irradiance	eruptions
1	+0.02	-0.24	0.00
2	+0.14	+0.27	+0.10
3	+0.32	-0.09	+0.03
Mean	+0.30	-0.04	+0.09

- By integrating proxy data with climate modelling, we can use past climatic changes to study the dynamics of the climate system.
- Orbital changes can explain the long-term upward trend in ENSO variability over the past 8 ka. Decreasing summer insolation results in a weakening of the Asian monsoon, reducing the stability of the tropical Pacific and making it easier for El Niño events to develop.
- On shorter timescales, there is no evidence that natural or anthropogenic forcings influence the amplitude of ENSO variability. This supports the notion that ENSO is a system where variability arises from internal dynamics, independent of external forcing.

References

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- Phipps, McGregor, Gergis, Gallant, Neukom, Stevenson, Ackerley, Brown, Fischer and van Ommen (2013), Paleoclimate data-model comparison and the role of climate forcings over the past 1500 years, *Journal of Climate*, doi:10.1175/JCLI-D-12-00108.1, in press.