

Quantifying changes in ENSO dynamics over the Holocene

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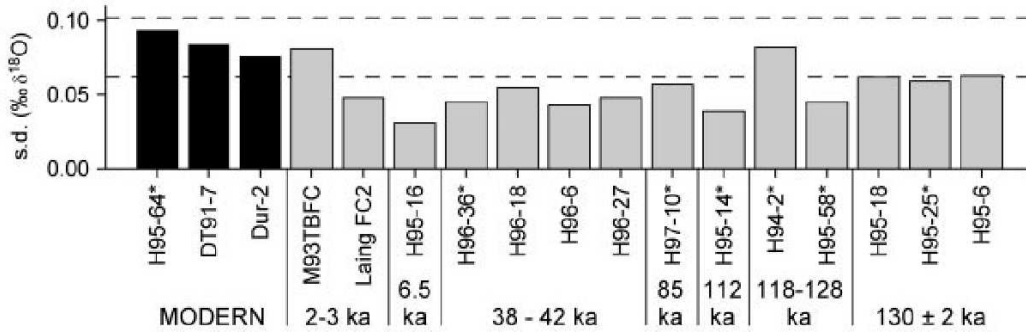
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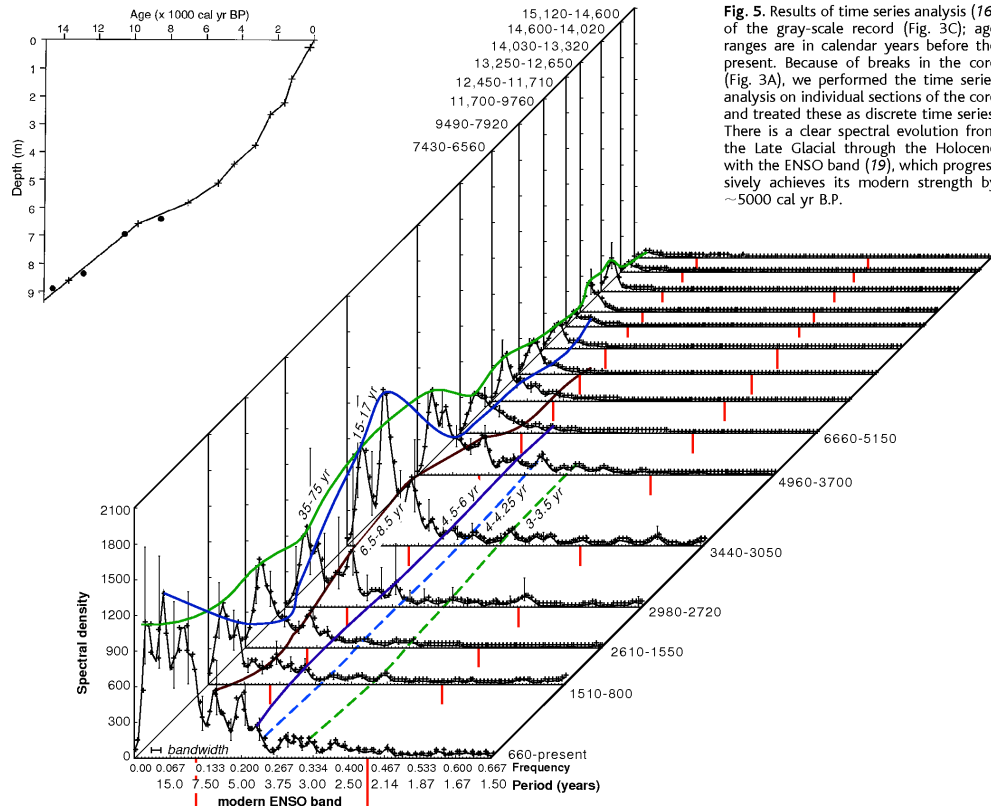
²Centre for Australian Weather and Climate Research, Hobart, Australia

El Niño has changed...

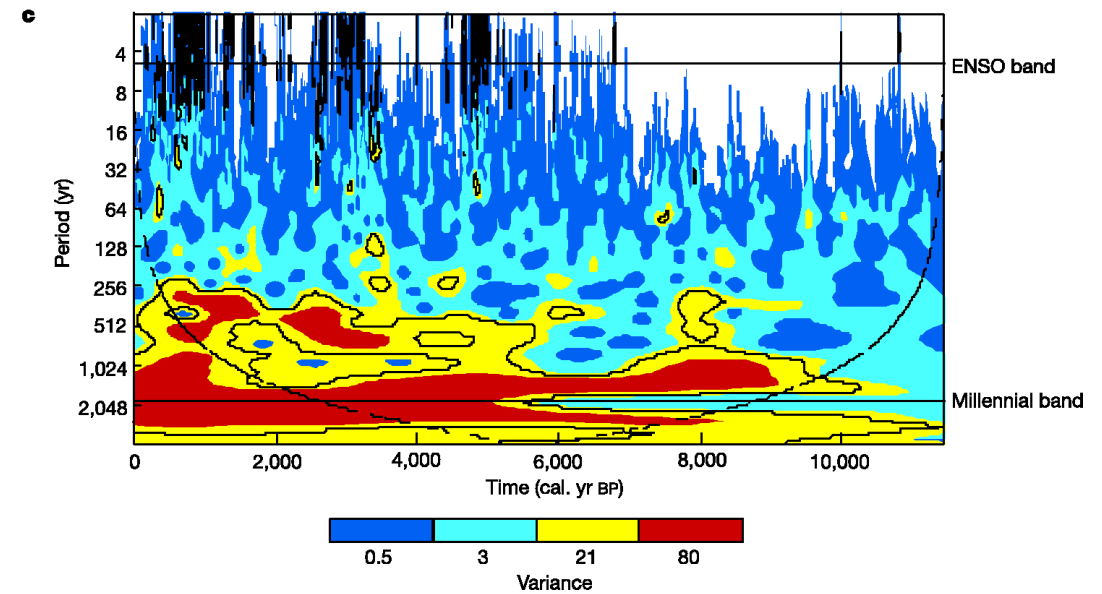
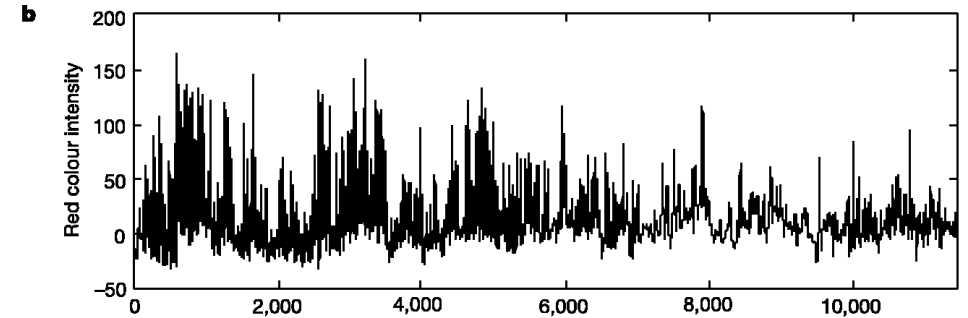
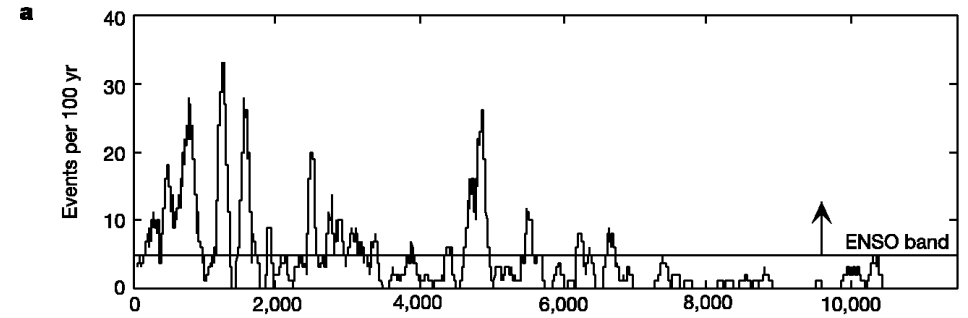
- Proxy reconstructions from across the Pacific Basin show that:
 - “Modern” El Niño began 7-5 ka BP, with only weak decadal-scale events beforehand
 - El Niño was 15-60% weaker at 6 ka BP than at present
 - Gradual strengthening of El Niño thereafter
 - Evidence of a peak in strength at 2-1 ka, possibly earlier in the western Pacific than in the east



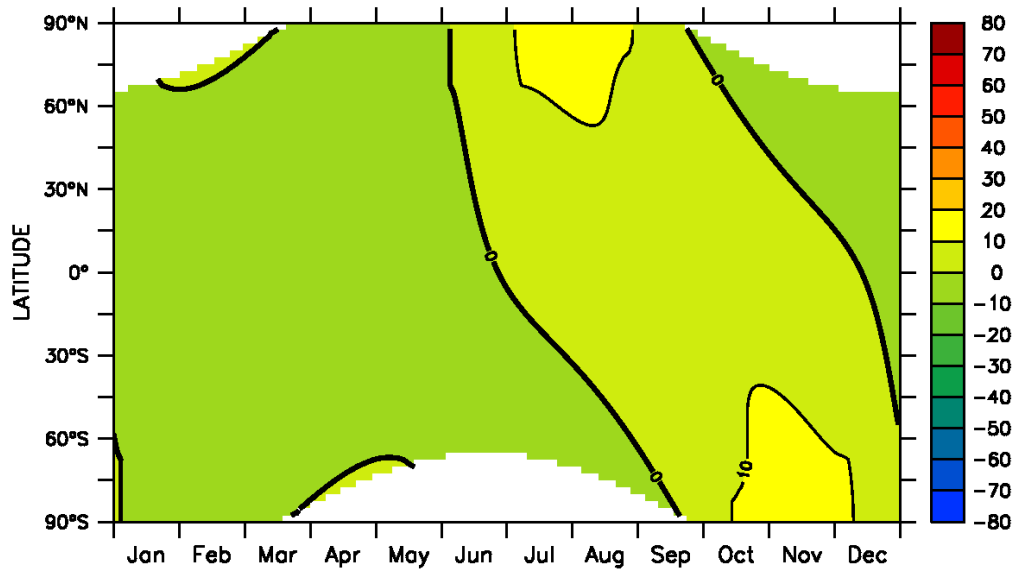
Tudhope et al. (2001), *Science*



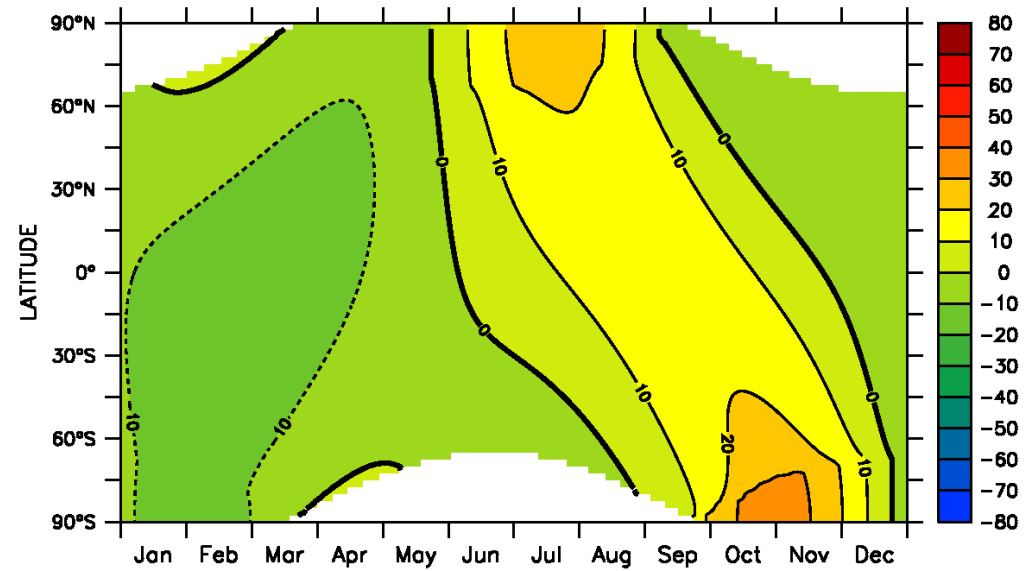
Rodbell et al. (1999), *Science*



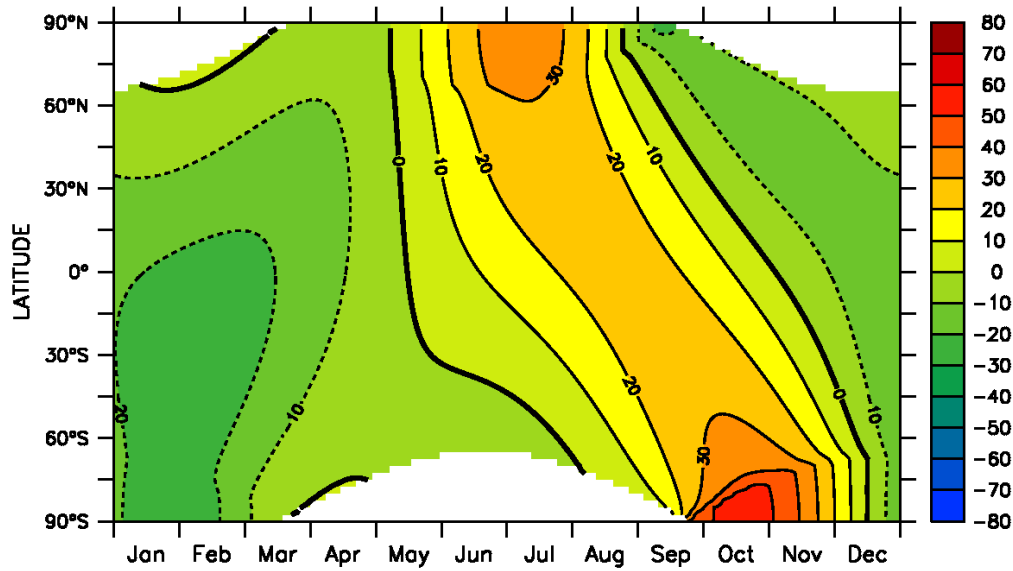
Moy et al. (2002), *Nature*



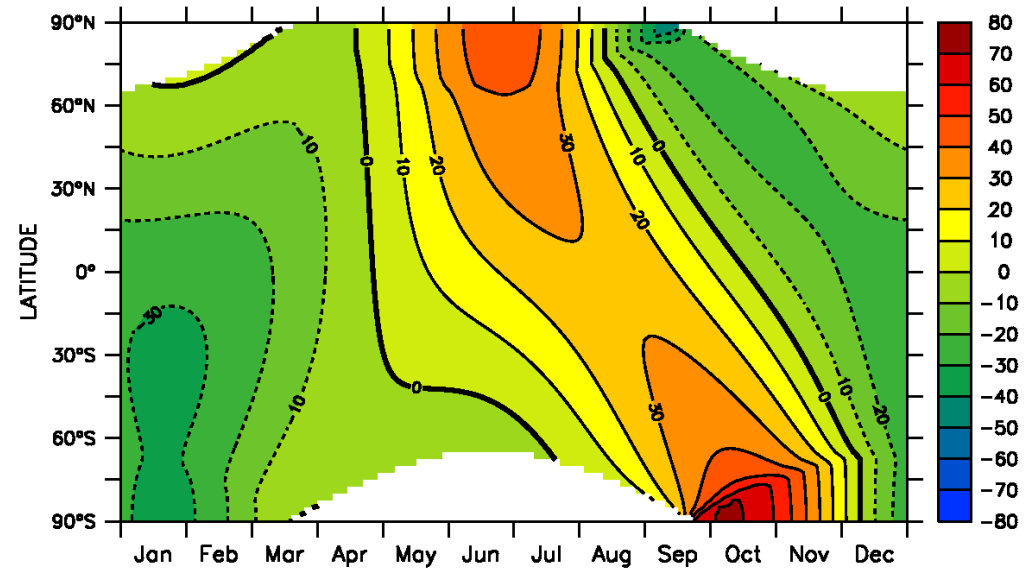
Insolation anomaly (Wm^{-2}): 2ka BP



Insolation anomaly (Wm^{-2}): 4ka BP



Insolation anomaly (Wm^{-2}): 6ka BP



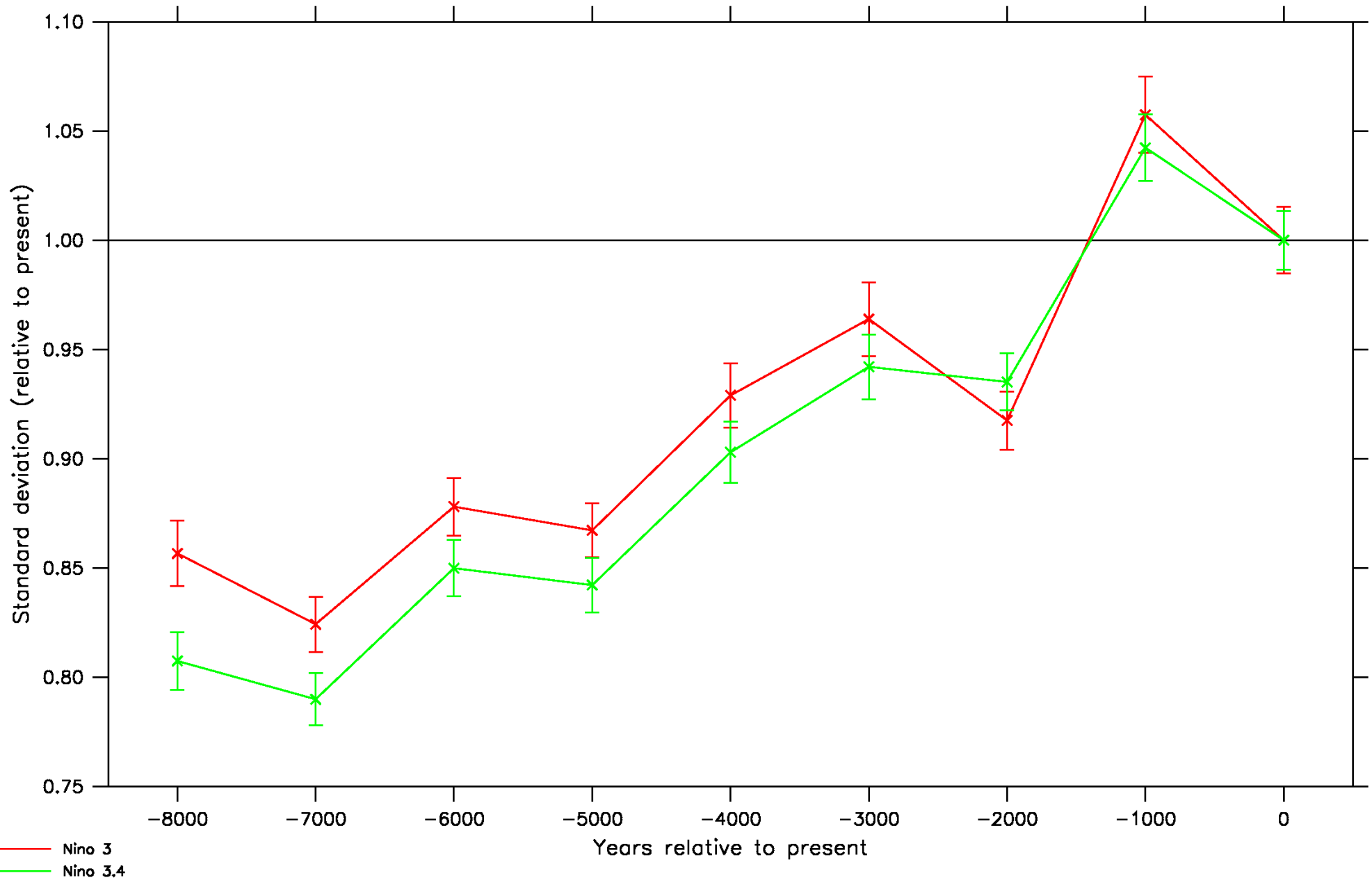
Insolation anomaly (Wm^{-2}): 8ka BP

Current understanding

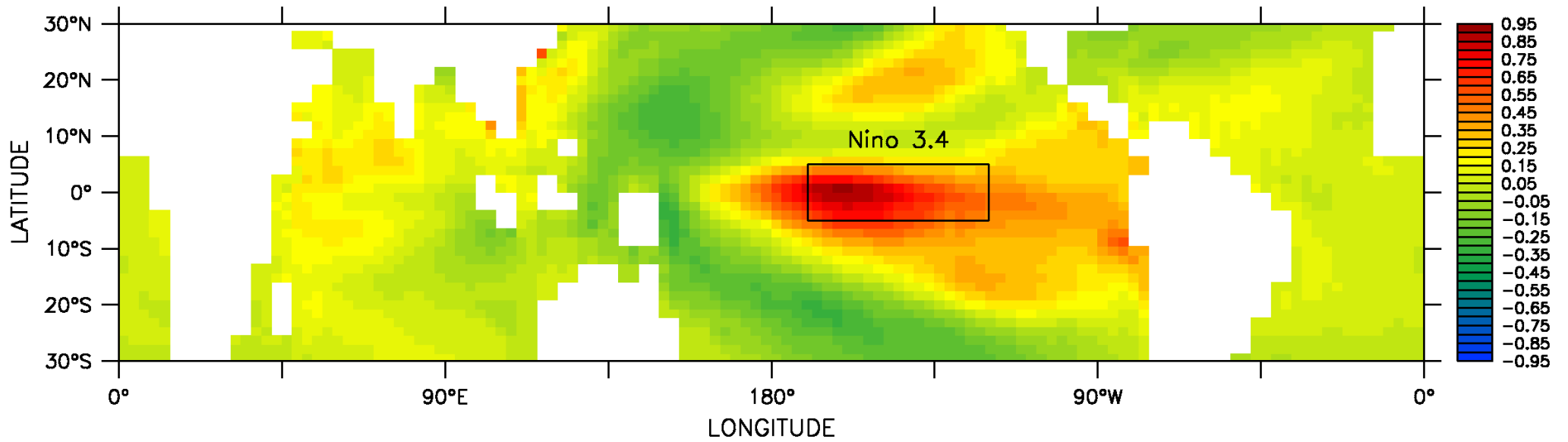
- Previous modelling work has shown that orbitally-driven changes in insolation can alter ENSO behaviour
- Broadly consistent mechanism found to explain weaker mid-Holocene ENSO:
 - Insolation changes result in enhanced seasonal cycle in NH
 - Intensification of summer monsoon system
 - Enhanced Walker circulation
 - Stronger easterly trade winds in central and western Pacific
 - Steeper thermocline/increased upwelling in central and eastern Pacific
 - Suppresses development of El Niño events
- However, this proposed mechanism is qualitative in nature and has yet to be rigorously tested

Simulations of the late Holocene climate

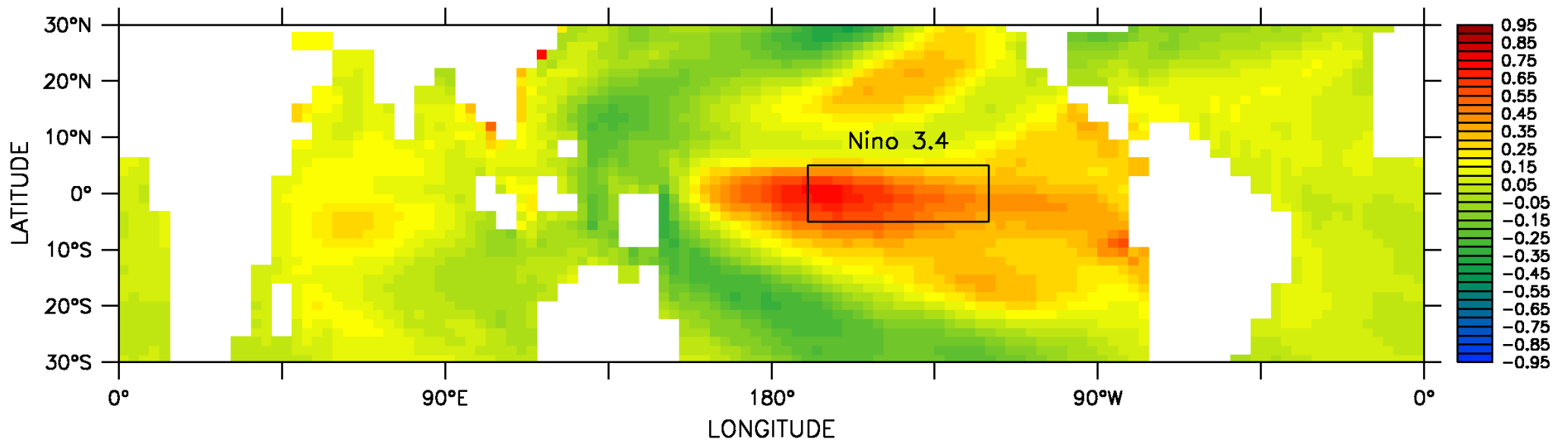
- CSIRO Mk3L climate system model v1.1:
 - Atmosphere: R21 ($5.6^\circ \times 3.2^\circ$), 18 vertical levels
 - Ocean: $2.8^\circ \times 1.6^\circ$, 21 vertical levels
 - Sea ice: Dynamic-thermodynamic
 - Land surface: Static vegetation
 - Flux adjustments applied
- Snapshot simulations for 8, 7, 6, 5, 4, 3, 2, 1 and 0 ka BP:
 - Only the Earth's orbital parameters are varied
 - Atmospheric CO₂ concentration = 280ppm
 - Solar constant = 1365 Wm^{-2}
 - Integrated for 1000 years



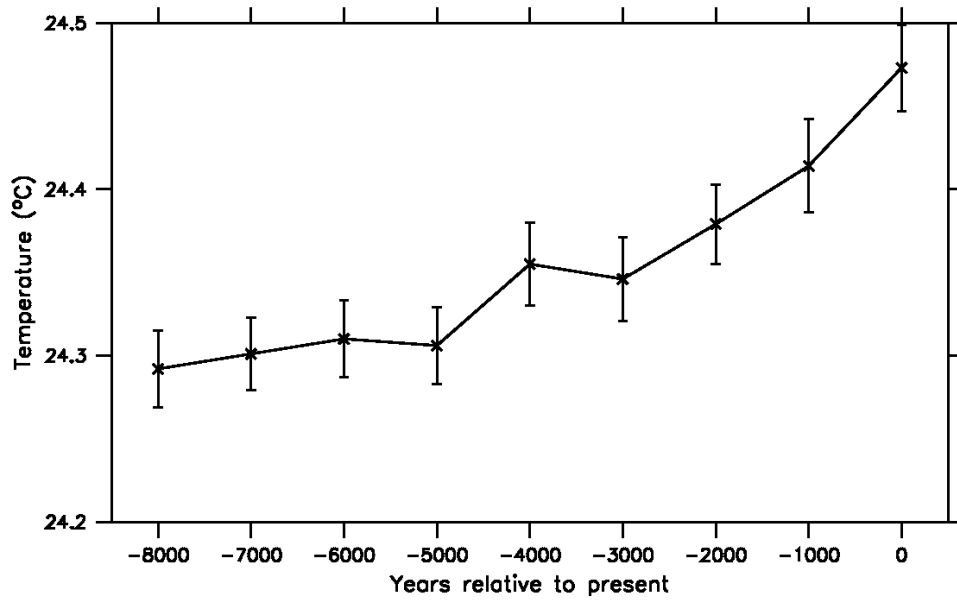
Standard deviation of Niño SST anomaly



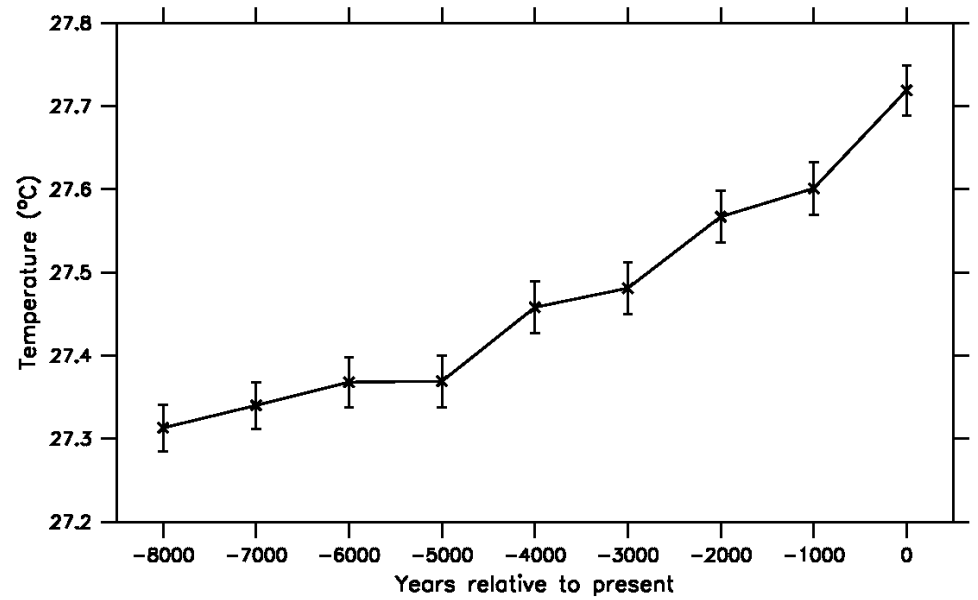
EOF1 of monthly SST anomalies (°C): 0ka BP



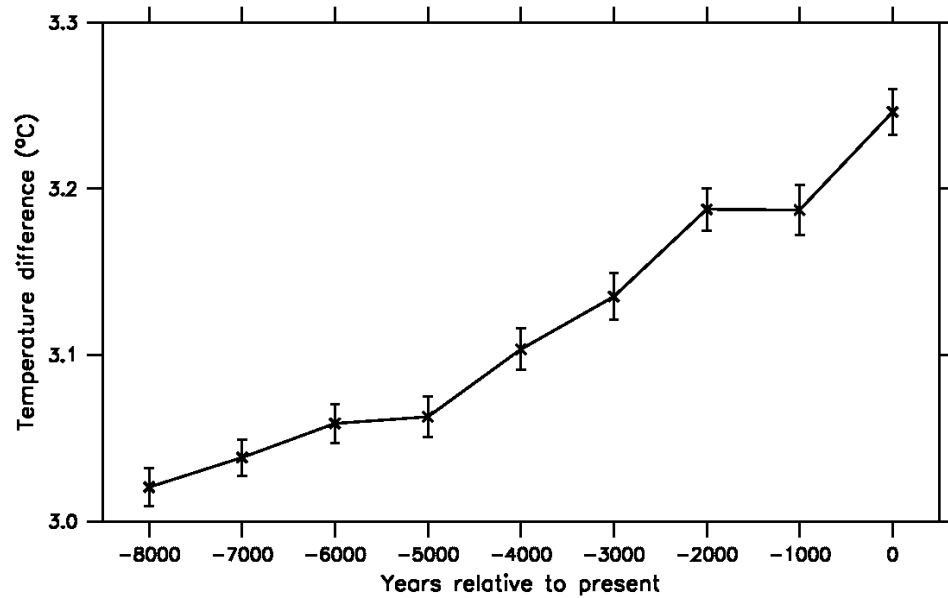
EOF1 of monthly SST anomalies (°C): 8ka BP



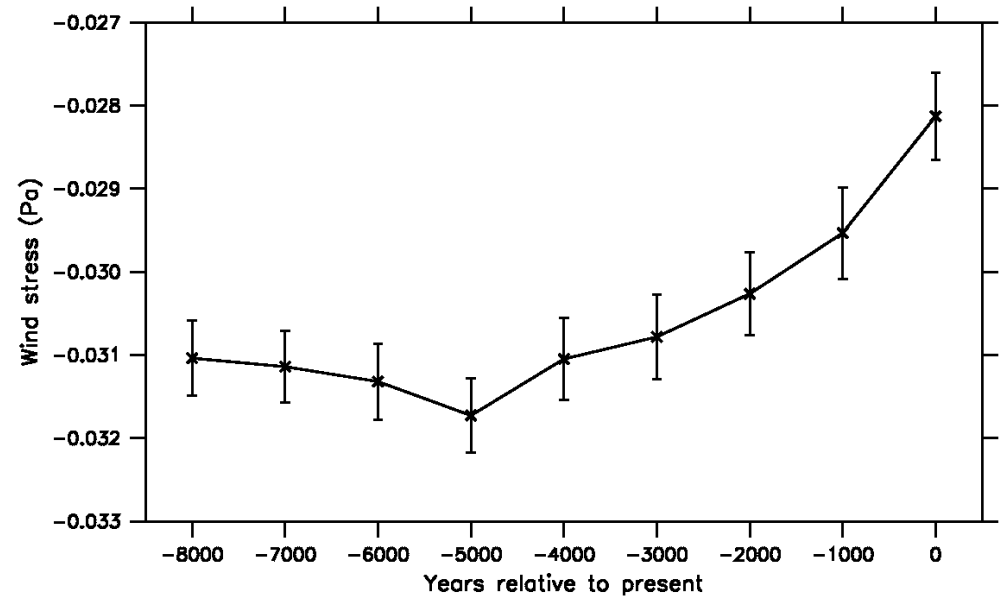
(a) Annual-mean Nino 3 SST



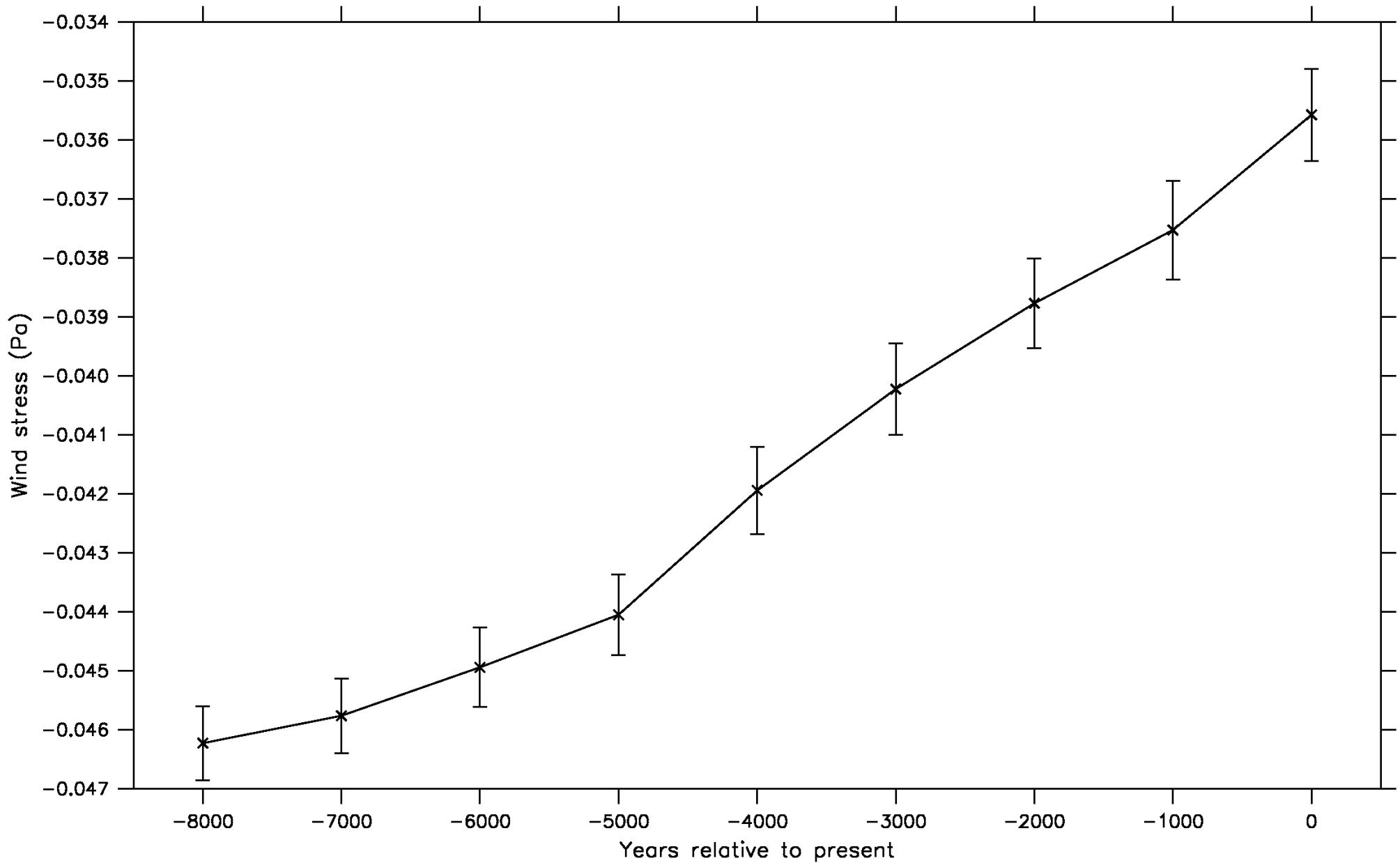
(b) Annual-mean Nino 4 SST



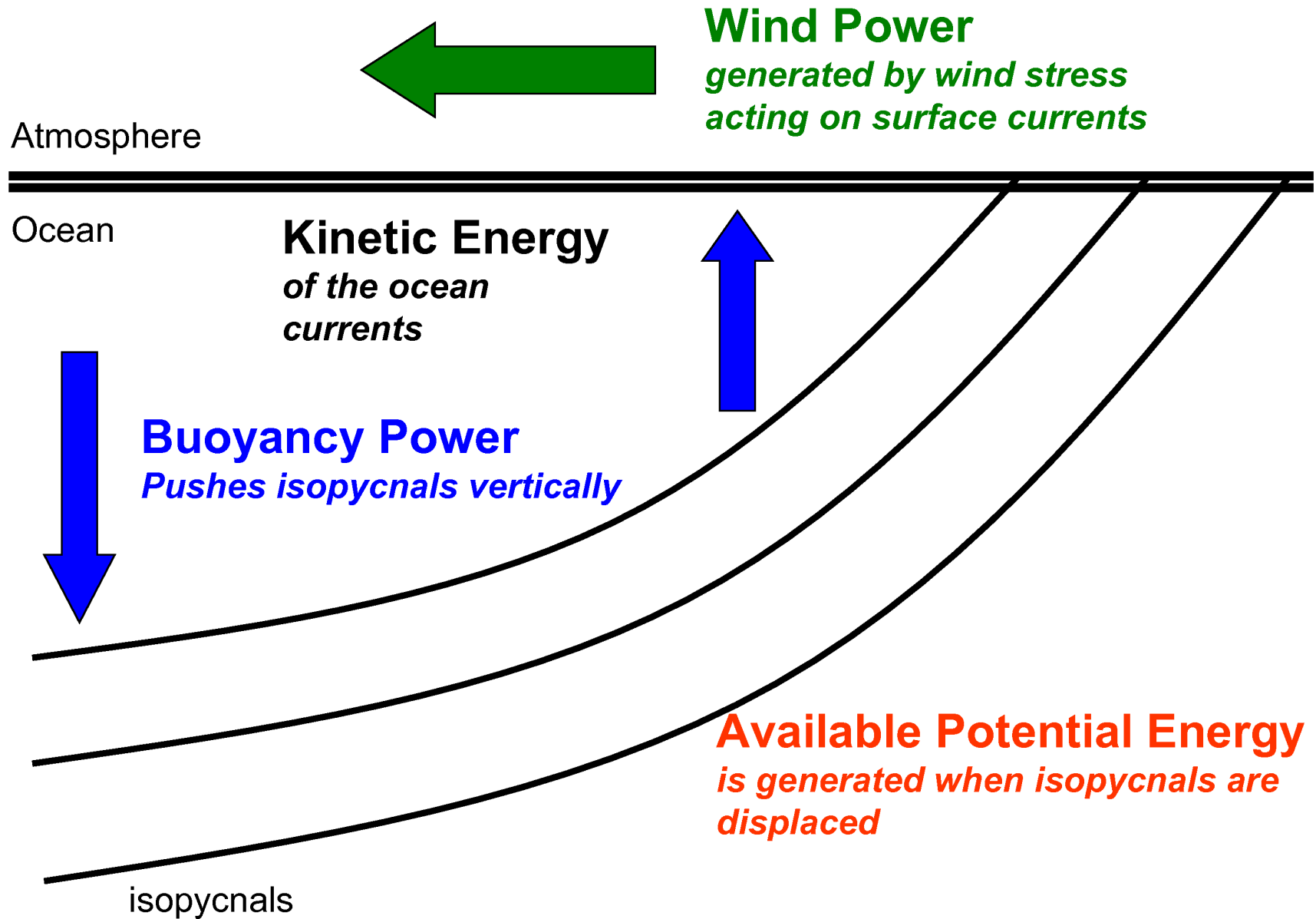
(c) Nino 4 SST minus Nino 3 SST



(d) Annual-mean Nino 4 zonal wind stress

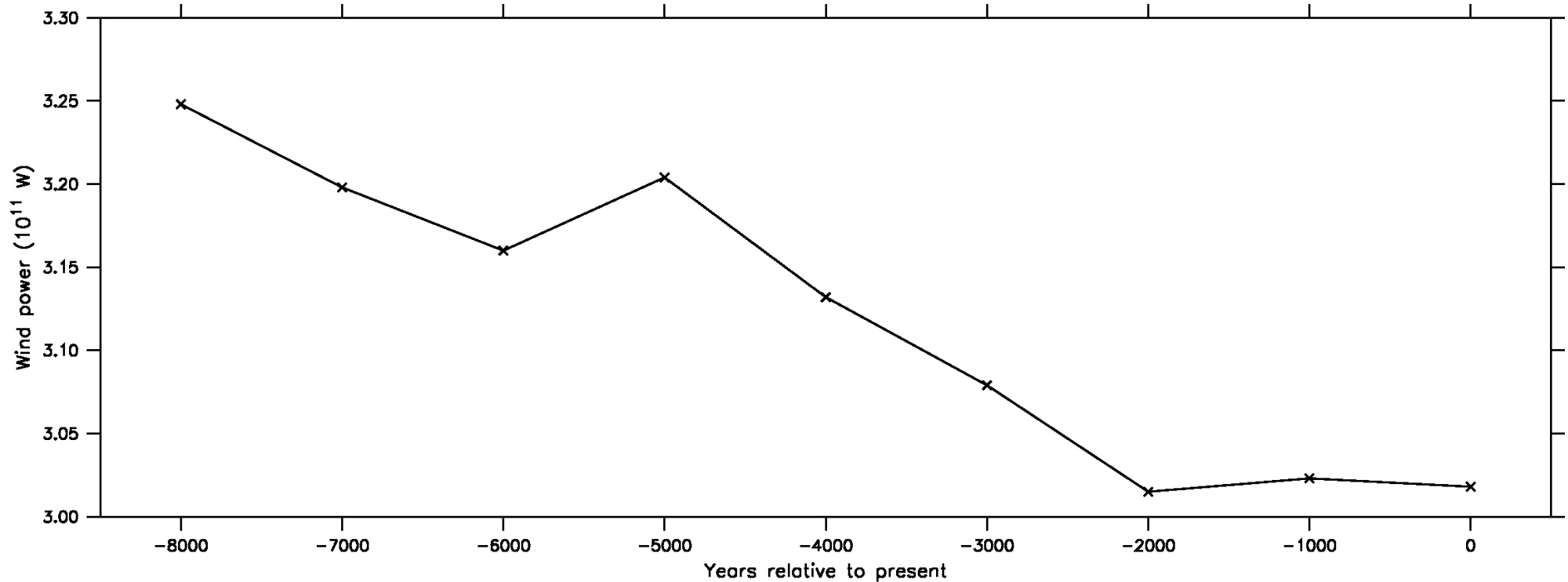


JASO zonal wind stress in Niño 4 region

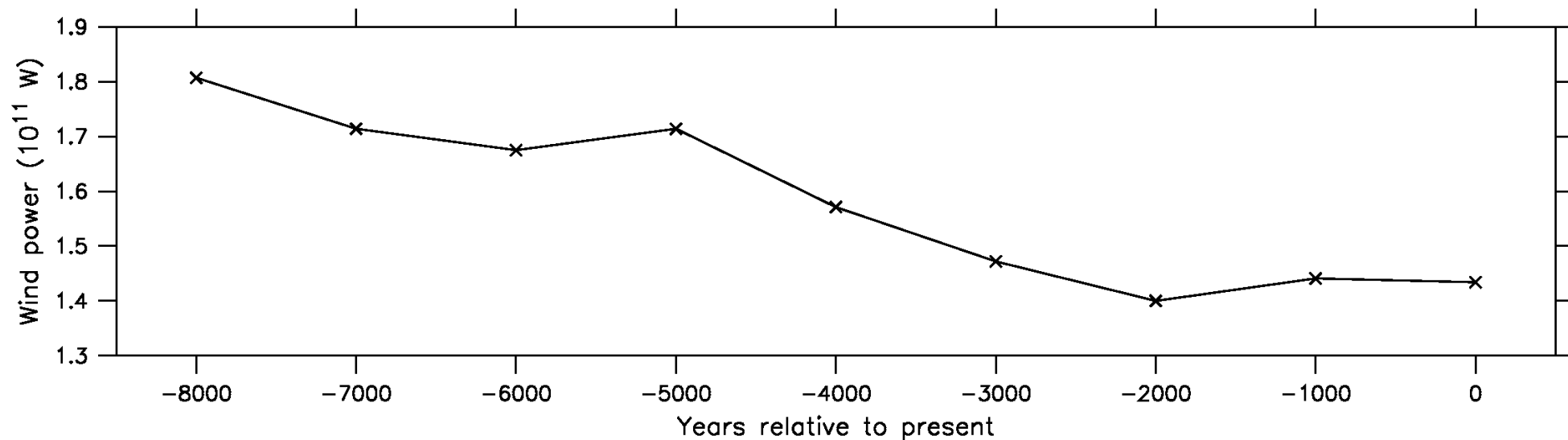


Wind power

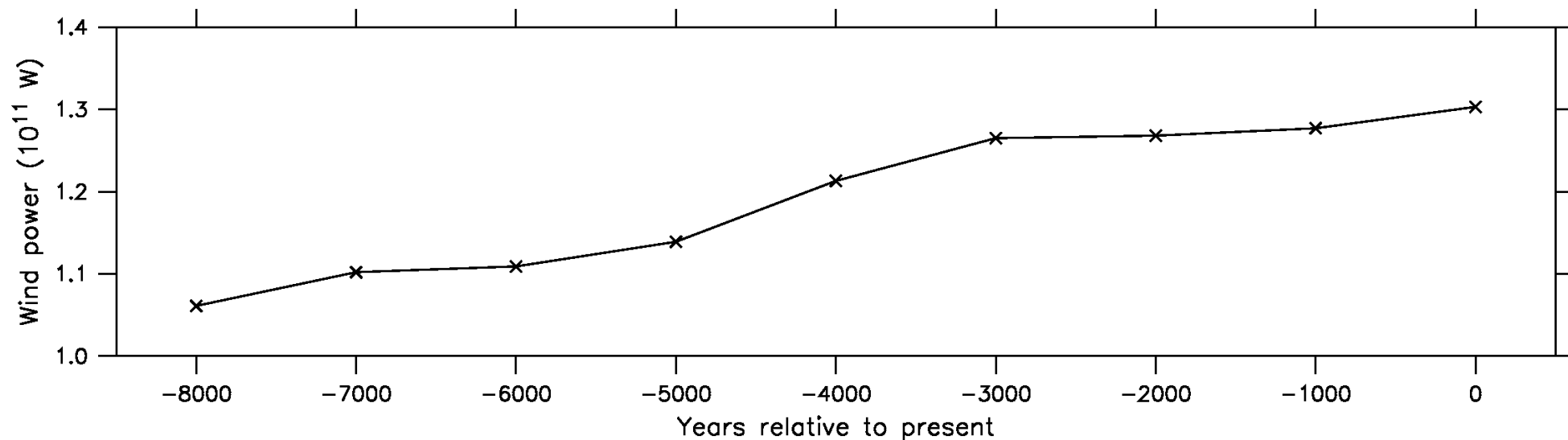
$$W = \iint_{z=0} \underline{u} \cdot \underline{\tau} \, dx dy$$



Annual-mean wind power

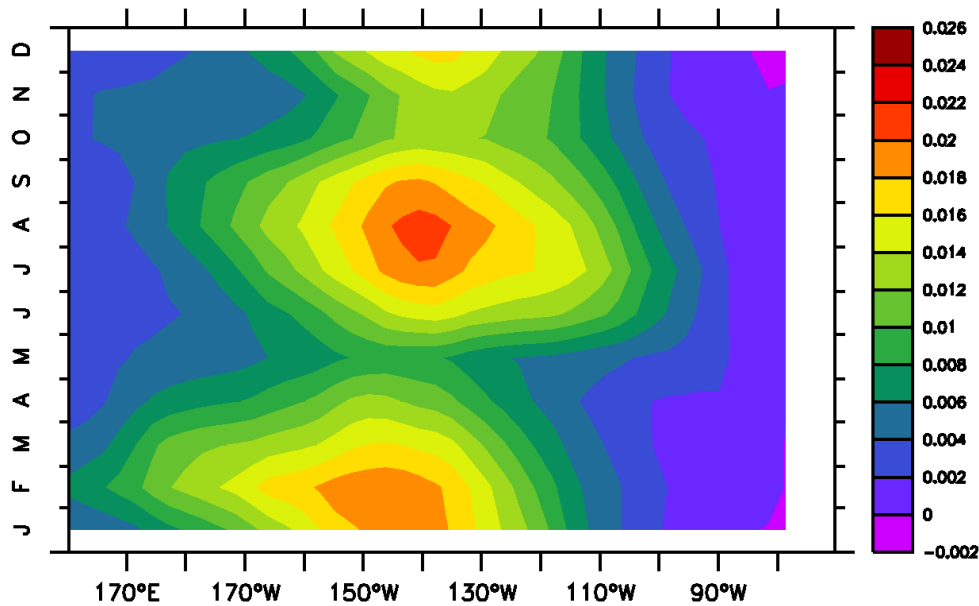


June–July–August–September wind power

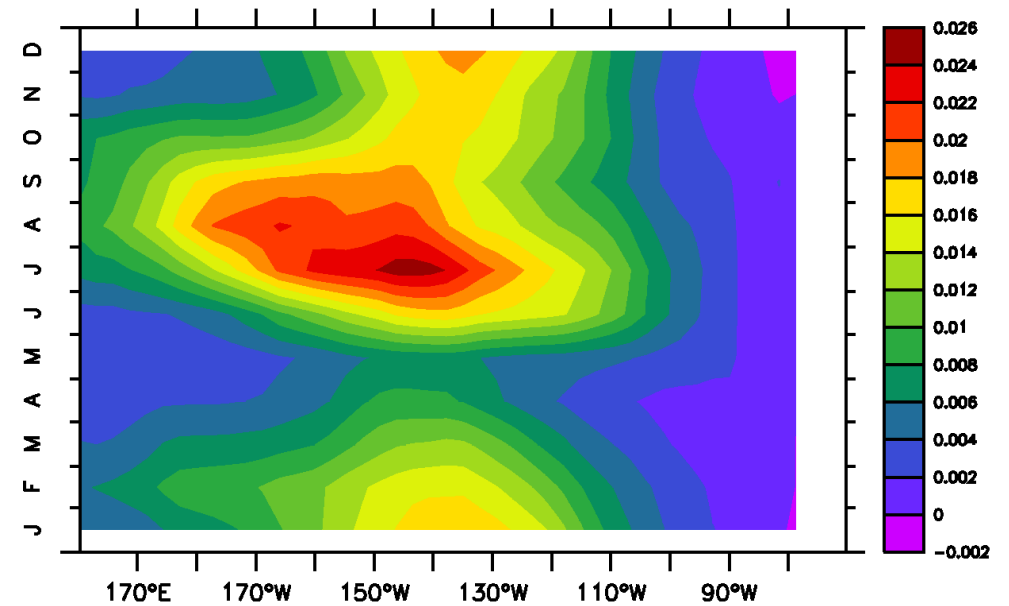


January–February–March–April wind power

Annual cycle

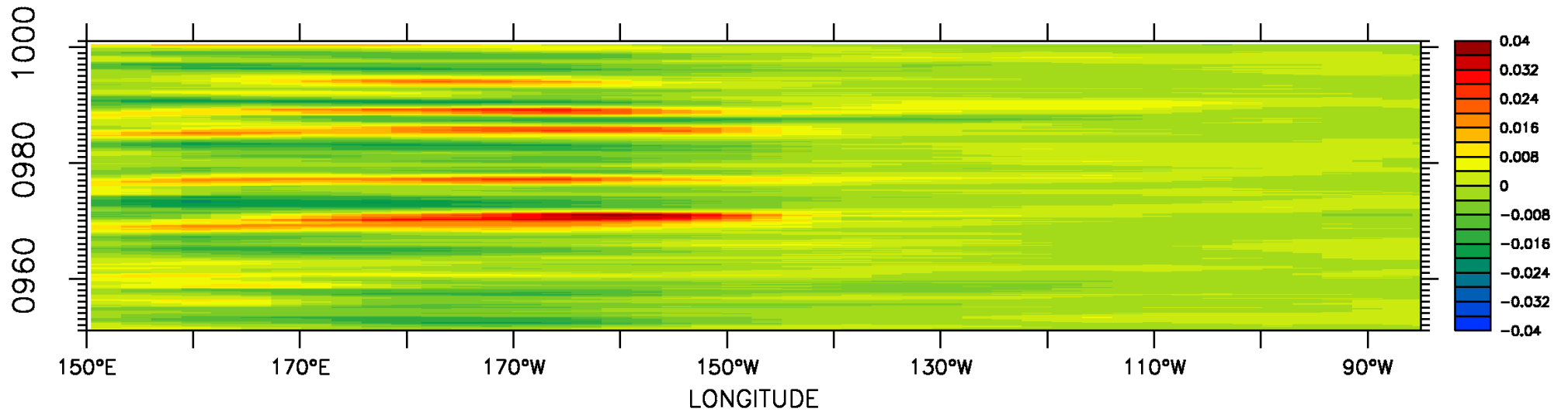


0 ka

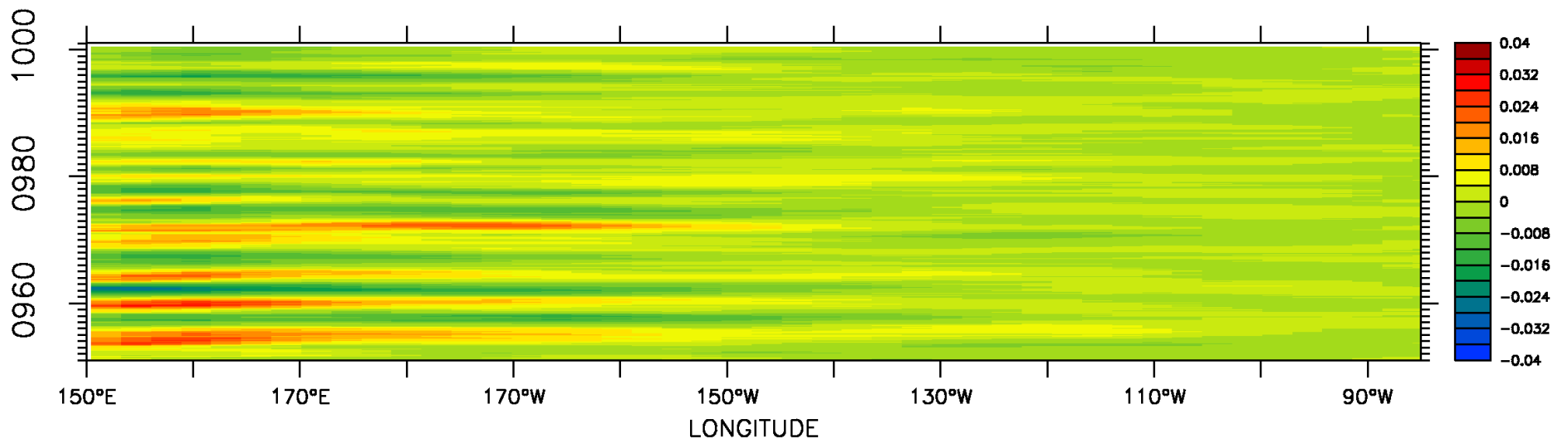


8 ka

Wind power (Wm^{-2})



Zonal wind stress anomaly at equator (Pa): 0ka BP



Zonal wind stress anomaly at equator (Pa): 8ka BP

Conclusions

- By forcing a model with orbitally-driven insolation changes only, we are able to broadly reproduce the changes in ENSO behaviour over the Holocene.
- Physical links between ENSO, the Walker Circulation and the Asian monsoon appear to explain the upward trend in variability.
- However, it does not explain the peak at 1 ka. Other mechanisms therefore appear to be at work.
- The key to understanding and *quantifying* past changes in ENSO behaviour may be to define better diagnostics.
- A full understanding of the processes that drive changes in ENSO variability may be within grasp. However, this will require an approach that integrates the theory, data and modelling communities.